

## Effects of soil water decline on diurnal and seasonal variations in sap flux density for differently aged Japanese cypress (*Chamaecyparis obtusa*) trees

K. Tsuruta, T. Kume, H. Komatsu, K. Otsuki

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**Abstract.** The effects of soil drought on transpiration are often neglected when predicting transpiration for forests in humid regions under the influence of the Asian monsoon. These effects have indeed been neglected for Japanese cypress, *Chamaecyparis obtusa*, a major plantation species in Japan and the surrounding area, probably because previous studies have reported no clear effects of soil drought on transpiration for Japanese cypress forests. However, a few studies have reported an apparent reduction in transpiration with soil drought for young Japanese cypress forests. It remains unclear whether such a reduction in transpiration is limited to young Japanese cypress forests or if it is not uncommon for mature Japanese cypress forests, which occupy a large area in Japan. To clarify this point, we conducted sap flux measurements in a year with soil drought on three differently aged Japanese cypress stands including mature (43 years old) and relatively young (23 and 26 years old) trees. In a diurnal time scale, a cross correlation analysis of sap flux density ( $F_d$ ) and vapor pressure deficit (VPD) showed that the time lags between  $F_d$  and VPD were 1-3 h in dry soil conditions. These were larger than those of wet soil conditions (<1 h) for all sample trees.  $F_d$  at a given VPD in dry soil conditions was smaller than that in wet soil conditions for all sample trees; a 28%–63% reduction in the rate of change in  $F_d$  was observed under dry soil conditions. Because our results were obtained when the non-exceedance probability of recorded monthly precipitation was 9%–18%, the results suggest the need to consider the effects of soil drought more extensively. Those effects should be considered for not only relatively young but also mature Japanese cypress when predicting diurnal and seasonal patterns of transpiration in years with soil drought, and when predicting inter-annual patterns of transpiration for Japanese cypress despite humid temperate climate.

**Keywords** *Chamaecyparis obtusa*, soil drought, sap flux density, transpiration, water deficit, tree age

**Authors.** Kenji Tsuruta (ktsuruta@ffpri.affrc.go.jp) - Soil Resources Laboratory, Department of Forest Soils, Forestry and Forest Products Institute, Matsunosato 1, Tsukuba, Ibaraki 305-8687, Japan; Tomonori Kume - School of Forestry and Resource Conservation, National Taiwan University, Taipei 106, Taiwan;

Hikaru Komatsu - Graduate School of Education, Kyoto University, Kyoto 606-8501, Japan; Kyoichi Otsuki - Kasuya Research Forest, Kyushu University, Sasaguri, Fukuoka 811-2415, Japan.

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## Introduction

Transpiration is a major component of water cycling in a forest ecosystem (Jasechko et al. 2013, Schlesinger & Jasechko 2014), and has strong coupling with carbon cycling as reflected in gross ecosystem production, tree growth, as well as in soil and root respiration (Law et al. 2002, Irvine et al. 2008, Ruehr et al. 2012). Examining transpiration is essential for accurately understanding and predicting water and carbon cycling in a forest ecosystem.

Canopy ( $G_c$ ) and aerodynamic conductance ( $G_a$ ) regulate transpiration (Monteith and Unsworth 1990).  $G_c$  is the more important factor controlling transpiration when compared with that of  $G_a$  for forests, particularly for those comprised of coniferous species (Jarvis and McNaughton 1986).  $G_c$  is controlled by environmental factors, such as vapor pressure deficit, solar radiation, and soil moisture (Jarvis 1976). However, the high soil moisture in forests of humid regions has often caused researchers to neglect the effects of soil drought on  $G_c$  when predicting transpiration (Ford and Vose 2007, Ewers et al. 2008, Kume et al. 2010). This is indeed the case for Japanese cypress plantation forests, a major forest type in Japan (Japan Forestry Agency 2012). Previous studies modelled  $G_c$  (or its substitutes, such as transpiration and evapotranspiration) for forests comprised of Japanese cypress (*Chamaecyparis obtusa*) or other tree species in Japan (Morikawa et al. 1986, Murakami et al. 2000, Tanaka et al. 2002, Watanabe et al. 2004, Kosugi et al. 2006, Kumagai et al. 2008, Komatsu et al. 2006, 2008, 2010, 2014).

Many previous studies have found no clear effects of soil drought on transpiration for

Japanese cypress forests (Kosugi et al. 2013, Sun et al. 2014, Tateishi et al. 2015, Tsuruta et al. 2015). However, Hattori et al. (1993) reported a reduction of daily-scale transpiration that was caused by soil drought during the growing season in a 12-year-old Japanese cypress forest with a canopy height of 7.0 m, using a micrometeorological approach. This phenomenon might be specific to young Japanese cypress forests because young or short trees have shallow root systems and therefore are easily affected by soil drought (Irvine et al. 2002, 2004; Kume et al. 2007). If this is the case, the models that ignore the effects of soil drought on transpiration should be useful for predicting transpiration in most cases in humid regions of Asia. It should be kept in mind that young Japanese cypress forests occupy only a limited portion of the total area of Japanese cypress forests in Japan (Japan Forestry Agency 2012). Also, at a diurnal time scale, the commencement and cessation of sap flux lag behind those of transpiration, and the time lag could be prolonged under soil drought conditions because of the effects of hydraulic resistance and capacitance (Phillips et al. 1997). The time lag between the occurrence of sap flux and transpiration might cause errors in transpiration estimates at a diurnal time scale (Phillips et al. 1999, 2004; Kumagai et al. 2009) and has implications related to modelling the soil-plant-atmosphere continuum.

Past attempts have been made to examine the effects of soil drought on transpiration for mature Japanese cypress forests but they have had some shortcomings. Many previous studies used the micrometeorological method that requires a scaffold tower (Hattori et al. 1993, Tanaka et al. 1996, Kosugi & Katsuyama

2007, Kosugi et al. 2007, 2013); however, this method is not easily applicable to mature forests with tall trees. Our objective was to examine whether the effects of soil drought on transpiration could be detected in mature and relatively young Japanese cypress forests. We measured transpiration recorded using the sap flux method before and during soil drought, a method that is easily applicable to both young, short and mature, tall forests.

## Materials and methods

### Study site

This study was conducted at the Geisha site of the Kasuya Research Forest, Kyushu University, Japan (33°38'N, 130°31'E; 50 m a.s.l.), located ~15 km east of Fukuoka City. Annual mean precipitation and air temperature observed from 2000–2004 were 1,560 mm and 16.2°C, respectively.

We conducted measurements in three adjacent Japanese cypress stands with different stand ages and tree sizes; stands were named small (S), medium (M), and large (L) in increasing order of stand age and tree size (Table 1). The shallow soil was less than 34 cm deep (Fig. S1).

Three sample trees were randomly selected from each stand (Table 1); all stands had sparse understories. The study period, from 22 August, 2006, to 17 November, 2006, experienced less precipitation in the late season (9 October to 17 November) when compared with the early season (22 August to 8 October).

### Observations

#### Environmental factors

Solar radiation, air temperature, relative humidity, and precipitation were measured at an open field located approximately 1 km northwest of the Geisha site. Solar radiation was measured

using a LI-200 pyranometer sensor (Li-Cor., Lincoln, NE, USA), and the air temperature and relative humidity were measured using a temperature/humidity probe (DTR503A, Vaisala, Helsinki, Finland). Data were scanned every 10 s and stored at 10-min intervals in a data logger (C10X, Campbell Scientific Inc., Logan, UT, USA). Precipitation was measured using a tipping bucket rain gauge (TK-1, Takeda Keiki Co., Tokyo, Japan). These data were also integrated every 10 min and then stored in the data logger (CR10X). Volumetric soil water content was measured near the sample trees at one point in each stand at a depth of 0–10 cm. The measurements were conducted using a dielectric aquameter sensor (EC-10, Decagon Devices, Inc., Pullman, WA, USA).

#### Sap flux measurement

Sap flux density was measured by the thermal dissipation method with Granier-type sensors (Granier 1987). Each sensor consisted of a pair of 20-mm-long probes, 2 mm in diameter. The probes were inserted into the sapwood approximately 150 mm apart, vertically. The upper probe, which includes a heater, was supplied with a constant 0.2 W of power. The heat was dissipated into the sapwood and vertical sap flux surrounding the probe. The temperature difference between the upper and lower probes was measured every 30 s and averaged every 30 min by a CR1000 data logger with an AM16/32 multiplexer (both from Campbell Scientific Inc.). The recorded temperature difference was converted to sap flux density ( $F_d$ ) according to Granier (1987).

After the sap flux measurements, we used a dye injection method to determine the hydro-active sapwood region in the xylem (Umebayashi et al. 2007, 2008; Tsuruta et al. 2010). Sapwood depth was determined by the stained region of the sample discs. Tsuruta et al. (2010) provides a more detailed description of the dye injection method. For all sample trees, sapwood thickness as determined by this

was approximately 2.0 cm (Table 1), which is the same length as the sap flow probe. Based on Granier's assumption that the Granier-type sensor integrated one section of the probe intercepting active  $F_d$  and the other section in contact with non-conducting xylem, we did not employ the correction proposed by Clearwater et al. (1999).

To consider temporal changes in circumferential variations of  $F_d$  within the xylem trunk, we measured  $F_d$  at two directions (north and south) within the xylem trunk. Previous studies reported that the circumferential variation in  $F_d$  would vary with different soil moisture conditions (Ueda & Yoshikawa 1994, Lu et al. 2000). We examined the relationship between circumferential variations in  $F_d$  and soil moisture. Relationships between  $F_d$  at the north and south sides were significantly different during the early and late seasons for four trees among nine trees (Fig. S2; analysis of covariance,  $p < 0.1$ ). We used mean  $F_d$  data of the north and south sides of the trunk in the following analysis.

### Method of analysis

Prior to analyzing the sap flux data, we characterized the intensity of soil drought in this study. The precipitation amount for the current study period was compared with that of

the long-term average. For this, non-exceedance probability of the  $i^{\text{th}}$  precipitation event ( $p_i$ ) in ascending order for  $N$  years was calculated as  $p_i = i/(N + 1)$  based on the Weibull plotting position equation (Shiiba et al. 2010). This calculation was made using precipitation data collected from 1936 to 2010 ( $N = 76$ ) at the Iizuka station of the Japan Meteorological Agency, located about 15 km east of the Geisha site. We calculated the  $p_i$  of monthly precipitation of September and October 2006 for 76 years. Monthly precipitation recorded in the Geisha site was strongly correlated to that of the Iizuka station (Shinohara et al. 2008). This suggests that monthly precipitation recorded in the Iizuka station can be substitute for that recorded in the Geisha site.

To examine whether soil drought affects transpiration on a diurnal time scale, we examined the time lag between  $F_d$  and the vapor pressure deficit (VPD) at an hourly time scale in both the early and late seasons to compare the differences in the time lags between the early and late seasons. Note that the time lag could be prolonged when soil drought affects transpiration (Phillips et al. 1997). To determine the time lag, a cross-correlation analysis was performed using a 1-h-interval dataset. We assumed the lag to be one that maximized the correlation coefficient (Phillips et al. 1999).

To examine whether the effects of soil

**Table 1** Characteristics of the three differently aged stands (named S (small), M (medium), and L (large) stands in increasing order of stand age and tree size) and sample trees

Stand	No.	Stand age (years)	DBH (cm)	Tree height (m)	Stem density (trees ha <sup>-1</sup> )	Sapwood depth (cm)	Sapwood area (cm <sup>2</sup> )	Leaf area <sup>1</sup> (m <sup>2</sup> )	Leaf area : Sapwood area (m <sup>2</sup> cm <sup>-2</sup> )
S	S1	23	12.6	9.2	2,600	1.7	57.0	24.8	0.43
	S2		8.4	7.7		2.3	43.9	12.0	0.27
	S3		9.3	7.0		2.2	49.0	7.7	0.16
M	M1	26	17.2	11.5	1,000	2.2	103.6	25.1	0.24
	M2		18.7	11.5		2.0	104.8	39.7	0.38
	M3		19.7	11.3		2.6	137.2	61.5	0.45
L	L1	43	34.0	17.3	n. a.	3.5	330.7	96.7	0.29
	L2		29.0	16.4		2.7	219.6	85.8	0.39
	L3		25.5	18.1		2.1	155.8	41.8	0.27

Note. <sup>1</sup>Leaf area was measured by destructive sampling (Tsuruta et al. 2010); DBH - diameter at breast height.

drought on transpiration were detected on a seasonal time scale, we calculated the reduction rate of  $F_d$  as follows (Pataki et al. 2000, Kume et al. 2007):

$$\text{Reduction} = 1 - \frac{f_{\text{dry}}(\text{VPD})}{f_{\text{wet}}(\text{VPD})} \quad (1)$$

where  $f_{\text{dry}}(\text{VPD})$  and  $f_{\text{wet}}(\text{VPD})$  are the fitted curves for the relationship between daytime mean  $F_d$  and daytime mean VPD in the early and late seasons, respectively.

In  $f_{\text{dry}}(\text{VPD})$  and  $f_{\text{wet}}(\text{VPD})$ , the regulation of transpiration by air temperature ( $T_a$ ) and  $R_s$  was not considered. Although we noticed that transpiration was affected by  $T_a$  as well as VPD and  $R_s$  (Jones 2014), a meta-analysis of transpiration data for five Japanese cypress forests (Komatsu et al. 2014) demonstrated that the effects of  $T_a$  on transpiration were minimal when  $T_a$  exceeds 18.8°C. This condition was generally satisfied during the study period (see the Results). Furthermore, transpiration for conifer species was mainly driven by VPD rather than  $R_s$  (Jarvis and McNaughton 1986), suggesting that VPD is a major driving force for  $F_d$  in sample trees.

## Results

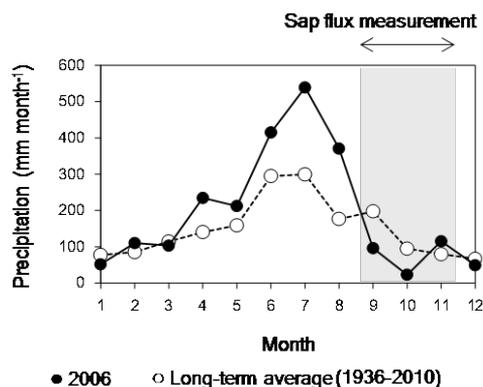
### Environmental conditions and sap flux density

Monthly precipitation in both September and October 2006 were smaller than the long-term average during the period from 1936 to 2010 (Figure 1). Monthly precipitation in September 2006 and the long-term average precipitation during September were 96 and 197 mm, respectively. Monthly precipitation of 96 mm in September was the 14<sup>th</sup>-smallest value ( $p_{14}$ ) for the 75 years (1936–2010). The  $p_{14}$  in September for the 75 years was 18.2%. Monthly precipitation in October 2006 and the long-term average precipitation during October

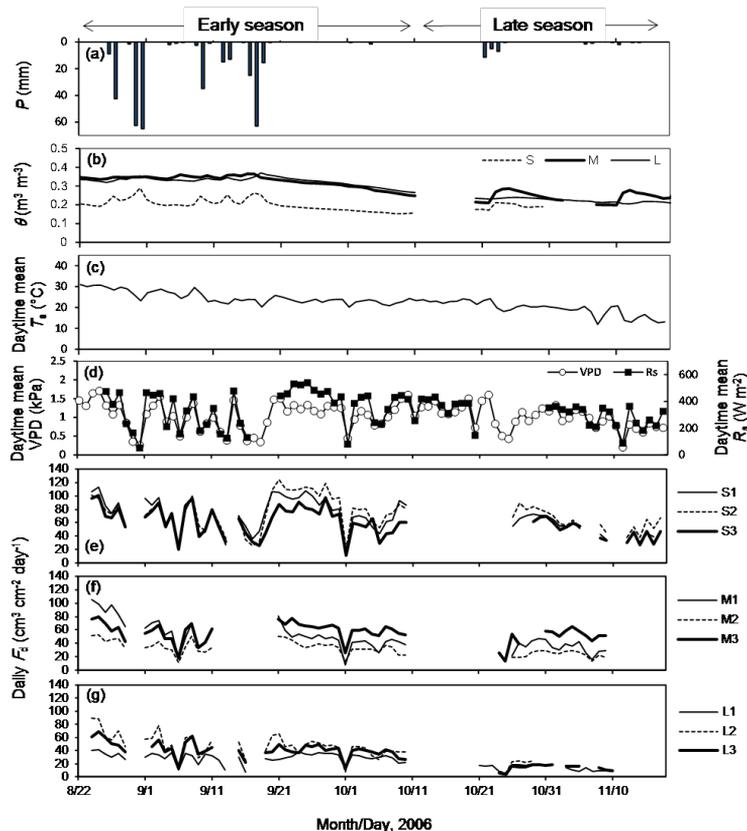
were 22.5 and 94.8 mm, respectively. Monthly precipitation of 22.5 mm in October was the 7<sup>th</sup>-smallest value ( $p_7$ ) for the 75 years. The  $p_7$  on October for the 75 years was 9.1%.

Precipitation occurred frequently in August and early September 2006 (Figure 2), but rarely occurred from late September onward. Soil water content continued to decline in the three stands after late September except after small precipitation events in late October and early November. The mean soil water contents in stands S, M, and L in the early season (22 August to 8 October) were 0.20, 0.33, and 0.33  $\text{m}^3 \text{m}^{-3}$ , respectively. The mean soil water contents in stands S, M, and L in the late season (9 October to 17 November) were 0.19, 0.24, and 0.22  $\text{m}^3 \text{m}^{-3}$ , respectively. In stand S, the mean soil water content in the late season was different from that in the early season only slightly, probably because data for the latter half of the late season were missing due to damage to probes.

VPD and  $R_s$  did not show clear seasonality during the study period (Figure 2). Day-to-day variations in daytime mean  $F_d$  corresponded to those of daytime mean VPD and  $R_s$  for all trees (Figure 2).  $F_d$  for stand S was larger than those for stands M and L.  $F_d$  in the late season was smaller than that in the early season for the three stands.



**Figure 1** Seasonal variations in long-term (1936–2010) average precipitation and that in 2006



**Figure 2** Day-to-day variations in (a) precipitation (P), (b) soil water content at depth of 10 cm ( $\theta$ ) for small (S), medium (M), and large (L) stands, (c) daytime mean air temperature ( $T_a$ ), (d) daytime mean vapor pressure deficit (VPD) and solar radiation ( $R_s$ ), sap flux density ( $F_d$ ) for stands (e) S, (f) M, and (g) L for the early and late seasons. A power outage and damage to probes caused some data to be lost.

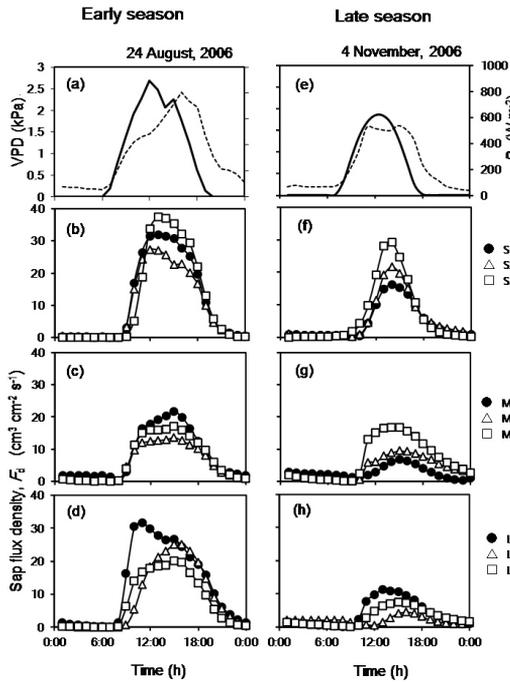
**Effects of soil drought on diurnal variations in sap flux density**

Figure 3 shows diurnal courses of VPD,  $R_s$ , and  $F_d$  of the trees in the three stands on fine days in the early (24 August 2006) and the late (4 November 2006) seasons, in which relatively large VPD,  $R_s$ , and  $F_d$  were observed.  $F_d$  for S trees was larger than those for M and L trees in both seasons. The peak values of  $F_d$  were smaller in the late season than in the early season for all three stands. This tendency was generally maintained when we used data

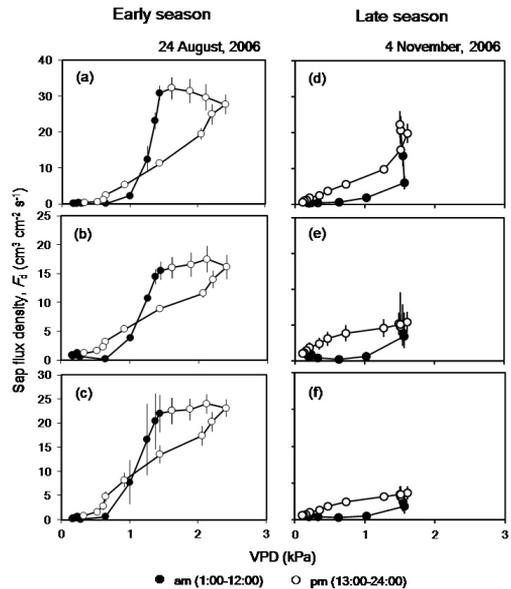
for other days.  $F_d$ -VPD relationships within a day were different between the early and late seasons (Figure 4). In the early season, diurnal patterns of  $F_d$  showed hysteresis with VPD in a clockwise manner for all three stands;  $F_d$  in the morning was larger than that in the evening at a given VPD. Rapid increases in  $F_d$  were observed in the morning. Relatively large  $F_d$  was maintained in the daytime (13:00–16:00). However, in the late season, the diurnal patterns of  $F_d$  showed hysteresis with VPD in a counter-clockwise manner for all three stands;  $F_d$  in the morning was smaller than that in the evening at a given VPD. Rapid increases in  $F_d$  were not observed in the morning.

The time lag between  $F_d$  and VPD was determined (Table 2) and was generally longer in the late season

than in the early season. In the early season, the mean time lag of stands S, M, and L was  $0, 0.3 \pm 1.2,$  and  $0.3 \pm 0.6$  h, respectively. In the late season, the mean time lag of stands S, M, and L was  $0.7 \pm 0.6, 1.7 \pm 0.6,$  and  $1.7 \pm 1.2$  h, respectively. The time lag between  $F_d$  and  $R_s$  was also generally shorter in the early season than in the late season. In the early season, the mean time lag of stands S, M, and L was  $1.3 \pm 0.6, 2.0 \pm 0,$  and  $1.7 \pm 0.6$  h, respectively. In the late season, the mean time lag of stands S,



**Figure 3** Representative diurnal trends of vapor pressure deficit (VPD, broken line), solar radiation ( $R_s$ , solid line), and sap flux density ( $F_d$ ) of small (S), medium (M), and large (L) stands in the (a-d) early and (e-h) late seasons. The figure presents data from 24 August 2006 for the early season and 4 November 2006 for the late season.



**Figure 4** Relationships between sap flux density ( $F_d$ ) and vapor pressure deficit (VPD) for sample trees of small (S), medium (M), and large (L) stands in the (a-c) early and (d-f) late seasons. The figure presents data from 24 August 2006 for the early season and 4 November 2006 for the late season. Closed and open circles represent the data for periods from 1:00 to 12:00 and from 13:00 to 24:00, respectively. Error bars show the standard error.

**Table 2** Summary for cross-correlation analysis between vapor pressure deficit (VPD), solar radiation ( $R_s$ ), and sap flux density ( $F_d$ ) for wet and dry soil conditions. A time lag ( $\Delta T$ ) between  $F_d$  and VPD or  $R_s$  that was observed to have the highest cross-correlation coefficient (R) in each condition was listed. Data of four days for the  $F_d$ -VPD pair and three days for the  $F_d$ - $R_s$  pair in each season were used for this calculation

Tree No.	VPD				$R_s$			
	Wet		Dry		Wet		Dry	
	$\Delta T$ (h)	R						
S1	0	0.94	1	0.97	1	0.97	2	0.98
S2	0	0.93	1	0.95	1	0.98	2	0.97
S3	0	0.94	0	0.96	2	0.98	2	0.97
M1	-1	0.92	2	0.97	2	0.96	3	0.96
M2	1	0.98	2	0.97	2	0.95	3	0.93
M3	1	0.98	1	0.99	2	0.95	2	0.94
L1	0	0.98	1	0.97	1	0.94	2	0.97
L2	1	0.97	3	0.67	2	0.97	3	0.83
L3	0	0.97	1	0.98	2	0.96	3	0.96

M, and L was  $2.0 \pm 0$ ,  $2.7 \pm 0.6$ , and  $2.7 \pm 0.6$  h, respectively.

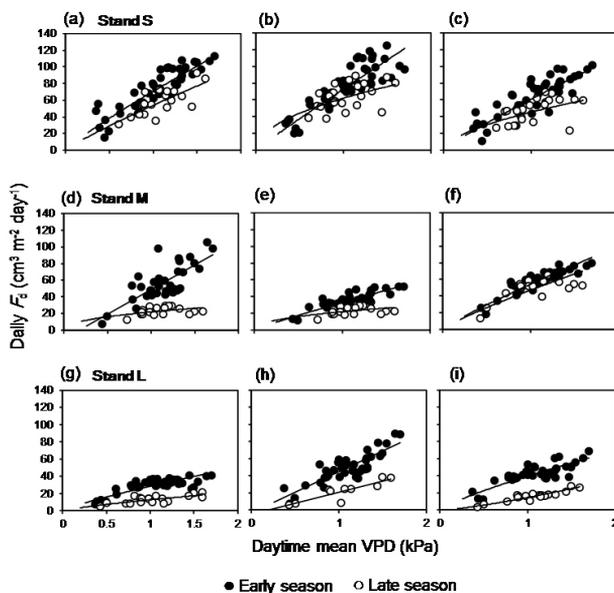
**Effects of soil drought on seasonal variations in sap flux density**

Daytime  $F_d$  at a given VPD was lower in the late season than in the early season for all three stands (Figure 5). Slopes of the  $F_d$ -VPD relationship in the early season were significantly different from those in the late season for S1–3, M4, 5, and L7 ( $p < 0.05$ ; analysis of covariance). Intercepts of the  $F_d$ -VPD relationship in the early season were significantly different from those in the late season for M6, L7, and L9 ( $p < 0.001$ ; analysis of covariance).

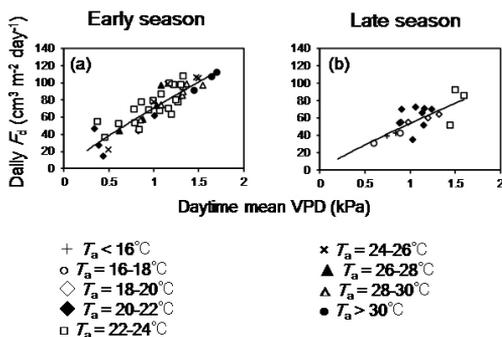
The difference in  $F_d$  between the late and early seasons did not seem to be caused by the difference in  $T_a$  between the seasons. Indeed, data recorded under different  $T_a$  classes were located along the regression line during both the early and late seasons (Figure 6; Figures S3, S4, S5). The data around 18.8°, which is the value of inflection temperature reported by Komatsu et al. (2014), were also located in the regression line.

The data around 18.8°, which is the value of inflection temperature reported by Komatsu et al. (2014), were also located in the regression line. Fig. 6 shows data only for a tree of stand S as an example. The reduction rates ranged from 0.19 to 0.33, 0.05 to 0.42, and 0.49 to 0.62 for

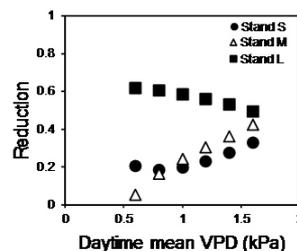
stands S, M, and L, respectively (Figure 7). The reduction rate was thus larger for stand L than for stands S and M throughout the range of VPD. A significant difference in a reference value of the reduction rate at daytime mean VPD = 1 kPa between stands S and L and be-



**Figure 5** Daily sap flux density ( $F_d$ ) in relation to daytime mean vapor pressure deficit (VPD) for small (S, a–c), medium (M, d–f), and large (L, g–i) stands for the early (closed circles) and late (open circles) seasons.



**Figure 6** Relationship between daytime mean vapor pressure deficit (VPD) and daily sap flux density ( $F_d$ ) for (a) the early and (b) late seasons for a tree of stand S. Data are classified according to air temperature ( $T_a$ ).



**Figure 7** Reduction rate in sap flux density ( $F_d$ ) for trees in small (S), medium (M), and large (L) stands. The reduction rate was averaged over three samples for each size.

tween stands M and L was observed ( $p < 0.05$ ; one-way analysis of variance).

## Discussion and conclusions

### Effects of soil drought on diurnal variations of sap flux density

Time lags observed in stands M and L were slightly larger than those for stand S in both the early and late seasons (Table 2). This could be associated with the differences in sapwood area among the different-sized trees (Table 1). Phillips et al. (1997) and Goldstein et al. (1998) reported that stem water storage capacity causing the time lag between transpiration and stem sap flow depended on the sapwood area which was related to tree size. Larger sapwood area might be a reason for the prolonged time lag observed in stands M and L.

The time lag between  $F_d$  and VPD was longer in the late season than in the early season (Table 2), and a counter-clockwise hysteresis was observed in the late season (Figure 4). These results suggest pronounced tree water recharge occurs during the evening to compensate for tree water discharge imposed by atmospheric demand in the daytime in the late season (Kumagai 2001, Phillips et al. 2003, Čermák et al. 2007, Hentschel et al. 2013).

### Differences in the effects of drought on transpiration at a seasonal time scale among trees with different sizes and ages

The reduction rate of  $F_d$  ranged from 0.05–0.62 in this study (Figure 7). Previous studies reported that Japanese cypress trees shed their leaves associated with air temperature from October to December (Tadaki & Kagawa 1968, Utsugi et al. 2001). We could not separate the effects of soil drought (leading to stomatal closure) and leaf fall on reduction of  $F_d$  in this study, because soil drought also accelerate leaf fall (e.g., Delzon & Loustau 2005).

However, Oren et al. (1999) reported that the leaf area reduction by a hurricane did not affect diurnal patterns of  $F_d$ , while the leaf area reduction affected absolute values of  $F_d$  for a *Taxodium distichum* forest. In this study, diurnal patterns of  $F_d$  in the late season were different from those in the early season (Figure 4, Table 2), indicating the detection of the effects of soil drought on  $F_d$  in this study.

We observed a pronounced decline of  $F_d$  caused by soil drought, when compared with previous studies in Japan. This might have occurred because previous studies were conducted during periods without severe soil droughts. For example, monthly precipitation during observations for the Kiryu Experimental Watershed was approximately the same as that of the long-term average (Kosugi et al. 2013). This suggests that the observation period lacked a period of severe soil drought. One exception reported by Hattori et al. (1993), which observed the effects of soil drought on transpiration in a Japanese cypress forest, reported the critical soil moisture at which transpiration started to decline. Transpiration for the Japanese cypress forest started to decline abruptly when the soil moisture passes the threshold value because of the limitation of extractable soil water. In this study, soil moisture may reach the critical value in the beginning of the late season. Kumagai et al. (2008) reported that soil moisture conditions did not have significant effects on  $F_d$  in both upper and lower slopes of a Japanese cedar (*Cryptomeria japonica*) forest, located about 60 km south from the present study site, although some dry spells occurred during the same period in that study as in the present study. Japanese cypress trees might have more characteristics that lead to water conservation than Japanese cedar trees under soil drought conditions (Nagakura et al. 2004).

In the present study, stand L had a higher reduction rate of  $F_d$  than stands S and M under soil drought conditions (Figure 7). This contrasts with our expectation. We expect-

ed smaller or no reduction in  $F_d$  for stand L under soil drought conditions because of the deeper root systems of these relatively large trees (Law et al. 2001; Irvine et al. 2002, 2004; Delzon et al. 2004; Kume et al. 2007). We do not have data related to root depth. However, the shallow soil (Figure S1) should not retain much soil water. The trees in the three stands also should not develop deeper root systems and access deeper soil water. The shallow soil might be one of reasons which all the trees in the three stands suffered from soil drought. One possible reason for the greater reduction in  $F_d$  for stand L might be a relatively large hydraulic resistance between the canopy and roots. A large hydraulic resistance may lead to higher stomatal sensitivity to dry conditions, when compared with relatively young and short trees (Hubbard et al. 1999, Phillips et al. 2002).

Differences in canopy structure and soil water dynamics may also have induced the larger reduction rate in  $F_d$  of the L trees. The L trees which have larger leaf area compared with the S and M trees may have relatively large wet-canopy evaporation. Moreover, whole-tree transpiration of L trees was much larger than those of the S and M trees (Tsuruta et al. 2008). Low sensitivity of soil water content to precipitation and the larger decrease of soil water content in Stand L might be caused by larger wet-canopy evaporation and root water uptake for stand L during periods of less precipitation in the late season. We need further investigations for clarifying the effects of tree size on  $F_d$ .

Our results were based on short-term measurements (about three months) and the number of replicates in this study was only three trees per one stand. However, these would not affect our conclusion that the reduction in transpiration was detected for not only S but also M and L stands. First, previous studies reported that seasonal trends in transpiration for temperate evergreen forests in Japan primarily responded to meteorological factors rather than

phenology (Komatsu et al. 2006, Kumagai et al. 2008, Komatsu et al. 2014). Note that trees within a stand are under approximately same meteorological conditions. Second, tree-to-tree variations in the reduction of transpiration with soil drought for each stand were relatively small (Figure 5). Third, tree-to-tree variations in transpiration seasonality are quite small for plantation forests in Japan (Komatsu et al. 2006). Based on these facts, the data of this study generally support our primary conclusion. Measuring physiological characteristics such as leaf water potential would be also useful for clarifying an underlying mechanism for the effect of soil drought on transpiration.

### Implications for modelling transpiration

Although previous studies have examined the effects of soil drought on transpiration for only relatively small and young Japanese cypress trees, we also observed the effects of soil drought on transpiration for mature Japanese cypress trees. Our results were obtained when the non-exceedance probability of recorded monthly precipitation during 75 years was 18.2% for September 2006, and 9.1% for October 2006, indicating that monthly precipitation on September and October less than that recorded on September and October 2006 could occur once per 5 and 11 years, respectively. This suggests that the effects of soil drought should be considered not only for young Japanese cypress trees, but also for mature ones when examining transpiration in years with soil drought and inter-annual variations in transpiration. Using transpiration models which neglect a function of soil moisture (Murakami et al. 2000, Tanaka et al. 2002, Watanabe et al. 2004, Kosugi et al. 2006, Ford and Vose 2007, Ewers et al. 2008, Kume et al. 2010, Sawano et al. 2015) for water resource management in Japan and surrounding regions could be problematic from this perspective. Our results suggest the need to include the effects of soil drought on transpiration in those models when

predicting transpiration from Japanese cypress forests despite humid temperate climate.

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## Supporting Information

The online version of the article includes Supporting Information:

**Fig. S1** Profile of the number of blows required for a penet-

tration of 10 cm ( $N$ ) in small, medium, and large stands

**Fig. S2** Relationship between daily sap flux density ( $F_d$ ) measured at north and south sides for small, medium, and large stands for the early and late seasons.

**Fig. S3** Relationship between daytime mean vapor pressure deficit (VPD) and daily sap flux density ( $F_d$ ) for the early and late seasons in the small stand.

**Fig. S4** Relationship between daytime mean vapor pressure deficit (VPD) and daily sap flux density ( $F_d$ ) for the early and late seasons in the medium stand.

**Fig. S5** Relationship between daytime mean vapor pressure deficit (VPD) and daily sap flux density ( $F_d$ ) for the early and late seasons in the large stand.